

# ANSWER KEY

Ocean 421  
Chemical Oceanography  
Spring 2000

Your Name \_\_\_\_\_

## Final Exam

(Use the back of the pages if necessary)(More than one answer may be correct.)

155  
+135  

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290 TOTAL POINTS

1. Due to the water molecule's ( $H_2O$ ) great abundance in seawater: (10 pts)
  - a) Seawater has a high heat capacity
  - b) The activity of  $H_2O$  is 1 in all equilibrium calculations
  - c) Salts dissolve easily in seawater
  - d) The activity coefficients of all dissolved constituents are equal
  - e) The activity of  $Ca^{2+}$  is greater than its concentration.
2. Due to the great abundance of dissolved salts in seawater: (10 pts)
  - a) sea water has a higher density than freshwater
  - b) sea water has a higher molar volume than freshwater
  - c) sea water has a higher boiling point than freshwater
  - d) sea water has a high ionic strength
  - e) sea water has a higher sound absorption than freshwater
3. Some elements in seawater are considered to be conservative.
  - a) What is the definition of a conservative element? (5 pts)

a non reactive element whose ratio to salinity (S) is constant in the ocean.
  - b) There are statements in the literature that uranium (U) is a conservative element in seawater. How would you design a study to show this is true or not true? (5 pts)

measure uranium together with a conservative element or property - such as salinity (S), potential temperature ( $\theta$ ) or  $Cl^-$ ,  $Ca^{2+}$  etc.

if U is conservative it will make a linear plot versus another conservative property

4. Both equilibrium and kinetic (dynamic) models have been proposed for the composition of seawater. What arguments can be given favoring the kinetic model relative to the equilibrium approach? (10 pts)

1. CLAYS COME FROM DETRITAL SOURCES RATHER THAN REVERSE WEATHERING
2. LACK OF EVIDENCE FOR SUFFICIENT MINERALS OF REVERSE WEATHERING ORIGIN
3. SW. COMPOSITION HAS CHANGED WITH TIME WHICH IS ALLOWED BY THE KINETIC MODEL BUT NOT THE EQUILIBRIUM MODEL.

5. As deep water flows from the North Atlantic to the North Pacific the carbonate alkalinity increases from 2350  $\mu\text{Eq}$  to 2475  $\mu\text{Eq}$  while total  $\text{CO}_2$  (e.g., DIC) increases from 2200  $\mu\text{M}$  to 2375  $\mu\text{M}$ . Estimate the average contributions of  $\text{CaCO}_3$  dissolution and organic carbon respiration that create this increase in DIC. (15 pts)

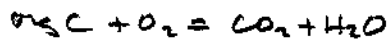
$$\Delta \text{ALIC} = 2475 - 2350 = 125 \mu\text{Eq} / 2 = 62.25$$

$$\Delta \text{DIC} = 2375 - 2200 = 175 \mu\text{M} - 62.25 = 112.75$$

ASSUME:



$$\Delta \text{DIC} = 1 \quad \Delta \text{ALIC} = 2$$



$$\Delta \text{DIC} = 1 \quad \Delta \text{ALIC} = 0$$

$$\frac{\Delta \text{DIC}_{\text{ORG}}}{\Delta \text{DIC}_{\text{CaCO}_3}} = \frac{112.75}{62.25} = 1.8$$

6.  $\text{CaCO}_3$  is added to water and it dissolves. How will the following parameters vary? (10 pts)

- pH ~~↓~~ ↑
- alkalinity ↑
- total  $\text{CO}_2$  ↑
- $\text{Ca}^{2+}$  ↑
- $P_{\text{CO}_2}$  ↓

7.  $^{238}\text{U}$  ( $t_{1/2} = 4.5 \times 10^9$  yr) decays to  $^{234}\text{Th}$  ( $t_{1/2} = 24.1$  d)

a) Define, and give the conditions for, secular equilibrium (5 pts)

DELAY ONLY  
 $t_{1/2P} \gg t_{1/2D}$

b) Assuming the activity of  $^{238}\text{U} = 2.3$  dpm  $\text{kg}^{-1}$  what is the molar concentration of  $^{234}\text{Th}$ ? (10pts)

$$A_{238\text{U}} = A_{234\text{Th}} = 2.3 \frac{\text{dpm}}{\text{kg}} = \lambda_{\text{Th}} N_{\text{Th}}$$

$$t_{1/2} = \frac{\ln 2}{\lambda}$$

$$\lambda = \frac{0.693}{24.1 \text{ d}} = 0.0288 \text{ d}^{-1}$$

$$N_{\text{Th}} = \frac{2.3 \text{ dpm kg}^{-1}}{0.0288 \text{ d}^{-1}}$$

c) Assume there has been an extreme bloom period with intensive scavenging that removes most of the  $^{234}\text{Th}$  from the water column in the euphotic zone. Suddenly the bloom stops. How long will it take for  $^{234}\text{Th}$  to reach secular equilibrium with  $^{238}\text{U}$ ? (5 pts)

5 half lives of daughter  
or  
 $5 \times 24.1 \text{ d} = 120 \text{ d}$

8. Which of the following is a first order process? (10 pts)

- a) River input
- b) Downwelling
- c) Biological Production, with respect to DIC
- d) Biological Production, with respect to N,P, Fe or another biolimiting element
- e) Radioactive decay

9. The Redfield (or RKR) ratio represents (10 pts)

- a) C:N:P content of organic matter formed during photosynthesis
- b) C:N:P content of nutrients released during respiration
- c) C:N:P of preformed nutrients in the surface ocean
- d) C:N:P of nutrient uptake during photosynthesis
- e) C:N:P of organic matter found in sediments

10. Relative to the Atlantic, Pacific deepwater (10 pts)

- a) pH is higher
- b)  $P_{CO_2}$  is higher
- c) Alkalinity is higher due to a larger organic C respiration:  $CaCO_3$  dissolution rate ratio
- d) Alkalinity is higher due to a smaller organic C respiration:  $CaCO_3$  dissolution rate ratio
- e) Total  $CO_2$  is higher than alkalinity

11. In which situations is it advantageous to use  $K'$ , the apparent equilibrium constant (10 pts)

- a) Ion activities (but not concentrations) can be measured and free energies are known
- b) Trace metal speciation calculations
- c) At ionic strength,  $I=0$
- d) Carbonate system equilibria in seawater
- e) When  $Q > K$ .

12. Due to the input of radionuclides during bomb testing in the 1960s, which information has been gained? (10 pts)

- a) Water mass ventilation using chlorofluorocarbon (CFC) tracers
- b) N-S mixing across the equator using  $^3H$  as a tracer
- c) Export production using  $C/^{234}Th$  ratios
- d) Stagnant boundary layer thickness using a  $^{14}C$  steady state box model
- e) The rate of turbulent mixing in the bottom boundary layer.

13. The oxidation of organic matter (10 pts)

- a) produces dissolved organic carbon
- b) produces dissolved inorganic carbon
- c) consumes energy
- d) provides bacteria with energy for growth and function
- e) is not possible in the absence of oxygen

14. Trace elements with nutrient-like distributions, such as barium and copper (10 pts)

- a) show a linear relationship with salinity and  $[Cl^-]$
- b) are depleted at the surface
- c) are enriched at the surface
- d) are regenerated (i.e. increase) with depth
- e) have fixed stoichiometric ratios to P

15. CO<sub>2</sub> and climate. There is evidence that the ocean's circulation varied in a regular way during the past glacial periods. This has to do with the different sources and strengths of bottom water formation.

- a) Broecker (1997) used a conservative geochemical tracer called PO<sub>4</sub>\* to determine the intensity of the northern and southern sources, where waters of southern origin are described by:

$$PO_4^* = PO_4 + O_2 / 175 - 1.95 \text{ mmol kg}^{-1}$$

O<sub>2</sub> ↓  
PO<sub>4</sub> ↑

and waters of northern origin by:

$$PO_4^* = PO_4 + O_2 / 175 - 0.73 \text{ mmol kg}^{-1}$$

Explain the origin of these equations and how it can be used for this purpose and why they are different for the northern and southern sources. (5 pts)

THE SUM OF  $PO_4 + \frac{O_2}{175}$  SHOULD REMAIN CONSTANT BECAUSE AS  $PO_4 \uparrow$   $O_2 \downarrow$  DUE TO RESPIRATION

THE LAST NUMBER IS THE PREFORMED  $PO_4$  WHICH IS DIFFERENT IN THE NORTHERN AND SOUTHERN SOURCES

WHY? PERHAPS BECAUSE THE SOUTHERN SOURCE IS UNDER THE ICE AND  $PO_4$  IS HIGHER BECAUSE THERE IS LESS BIOLOGICAL REMOVAL

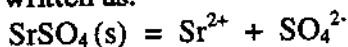
- b) What were the reorganizations in ocean circulation that occurred in the past and what was their origin? (5 pts)

NADW WAS SHUT DOWN AT SEMI-REGULAR INTERVALS DUE TO FRESHENING OF THE N. ATLANTIC SURFACE WATER

- c) If there are ocean reorganizations in the future, will they be caused by the same mechanism? If not what will be the cause? (5 pts)

AS  $CO_2 \uparrow$  SST  $\uparrow$  SO NADW MAY BE SHUT OFF BUT BECAUSE OF WARMING RATHER THAN FRESHENING.

16. Acantharian protozoa use strontium to form celestite,  $\text{SrSO}_4(\text{s})$  skeletons. Is the formation of celestite thermodynamically favorable in seawater? The solubility of celestite is written as:



- a) Calculate the equilibrium constant for this reaction at  $25^\circ\text{C}$  given the following free energies of formation ( $\Delta G_f^\circ$ ) and using  $(\Delta G_r^\circ) = -2.3RT \log K = -5.708 \log K$  at  $25^\circ\text{C}$ . (10 pts)

Species	$\Delta G_f^\circ$ (kJ mol <sup>-1</sup> )
$\text{SrSO}_4(\text{s})$	-1340.9
$\text{Sr}^{2+}(\text{aq})$	-563.83
$\text{SO}_4^{2-}(\text{aq})$	-744.53

$$\begin{aligned} \Delta G_r^\circ &= -563.83 + -744.53 - -1340.9 \\ &= -1308.36 + 1340.9 \\ &= +32.54 \text{ kJ mol}^{-1} \end{aligned}$$

$$\ln K = \frac{\Delta G_r^\circ}{-5.708} = \frac{32.54}{-5.708} = -5.70$$

$$K = 10^{-5.70}$$

- b) What is the equilibrium expression? (5 pts)

$$K = \frac{(\text{Sr}^{2+})(\text{SO}_4^{2-})}{(\text{SrSO}_4(\text{s}))}$$

1

- c) If seawater has  $[\text{Sr}^{2+}]_T = 0.0928$  mmol/L and  $[\text{SO}_4^{2-}] = 28.93$  mmol/L and the total activity coefficients are  $\gamma_{\text{Sr}} = 0.25$  and  $\gamma_{\text{SO}_4} = 0.065$ , what is  $\Delta G_r$ ? Is the formation of celestite thermodynamically favorable in seawater? (10 pts)

$$\begin{aligned} Q &= (\text{Sr}^{2+})(\text{SO}_4^{2-}) \\ &= [\text{Sr}^{2+}] \gamma_{\text{Sr},T} [\text{SO}_4^{2-}] \gamma_{\text{SO}_4,T} \\ &= (9.28 \times 10^{-5} \text{ M})(0.25)(28.93 \times 10^{-3})(0.065) \\ &= 4.36 \times 10^{-8} \\ &= 10^{-7.36} \end{aligned}$$

$$\begin{aligned} \Delta G_r &= \Delta G_r^\circ + 2.3RT \log Q \\ &= +32.54 + (5.708)(-7.36) \\ &= 32.54 - 42.01 \\ &= -9.47 \text{ kJ mol}^{-1} \end{aligned}$$

RXN TO RIGHT

17. A recent paper reported data for the gas nitrous oxide ( $N_2O$ ) in the surface waters of the Arabian Sea (Lal and Patra, 1998, Global Biogeochemical Cycles, 12, 321-327).

The average partial pressure of  $N_2O$  in the atmosphere over the Arabian Sea was 313 ppbv or  $10^{-6.50}$  atm. The Henry's Law constant for  $N_2O$  solubility at  $25^\circ C$  is  $K_H = 10^{-1.59} \text{ mol l}^{-1} \text{ atm}^{-1}$ .  $313 \times 10^{-9}$   
 $3.13 \times 10^{-7}$

a. What was the mean saturated concentration of  $N_2O$  in surface water in  $\text{mol l}^{-1}$ . (5 points)

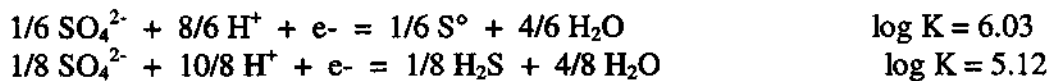
$$\begin{aligned} [N_{20}] &= K_H P_{N_{20}} = 10^{-1.59} \cdot 10^{-6.50} \\ &= 10^{-8.09} \\ &= 10^{0.91} \times 10^{-9} \\ &= 8.12 \times 10^{-9} \text{ M} \end{aligned}$$

b. The average degree of supersaturation was 130% and the average piston velocity for the average wind speed was  $22.7 \text{ cm hr}^{-1}$ . Calculate the average gas exchange flux in  $\text{mol cm}^{-2} \text{ d}^{-1}$  using the stagnant boundary layer model. (10 points)

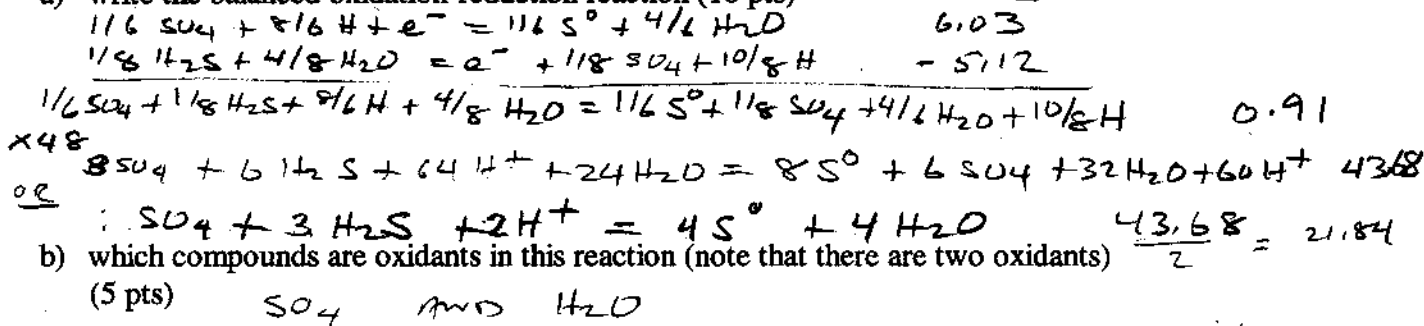
$$\begin{aligned} \text{Flux} &= \frac{D}{Z_{\text{FILM}}} (C - C_{\text{SAT}}) \\ &= \frac{D}{Z_{\text{FILM}}} (1.30 \times 10^{-8.09} - 10^{-8.09}) \\ &= 22.7 \frac{\text{cm}}{\text{hr}} (10.57 \times 10^{-9} - 8.12 \times 10^{-9}) \\ &= 22.7 \frac{\text{cm}}{\text{hr}} (2.37 \times 10^{-9} \frac{\text{mol}}{\text{l}}) \frac{\text{l}}{10^3 \text{ cm}^3} \\ &= \left( 22.7 \frac{\text{cm}}{\text{hr}} \right) \left( 2.37 \times 10^{-12} \frac{\text{mol}}{\text{cm}^3} \right) \\ &= 53.7 \times 10^{-12} \frac{\text{mol}}{\text{cm}^2 \text{ hr}} \times 24 \frac{\text{hr}}{\text{d}} \\ &= 1289 \times 10^{-12} \frac{\text{mol}}{\text{cm}^2 \text{ d}} \\ &= 1.289 \times 10^{-9} \frac{\text{mol}}{\text{cm}^2 \text{ d}} \end{aligned}$$

18. Sometimes different oxidation states of the same element can be both oxidants and reductants. In one case the oxidized form of sulfur ( $\text{SO}_4^{2-}$ ) can react with the reduced form ( $\text{H}_2\text{S}$ ) to form solid elemental sulfur ( $\text{S}^0$ ). Such a reaction may occur at the oxic-anoxic interface in the Black Sea.

Here are two relevant half reactions:



a) write the balanced oxidation-reduction reaction (10 pts)



b) which compounds are oxidants in this reaction (note that there are two oxidants) (5 pts)

$\text{SO}_4$  AND  $\text{H}_2\text{O}$

c) what is the equilibrium constant for this reaction (5 pts)

$$K = \frac{(\text{S}^0)^8 (\text{H}_2\text{O})^8}{(\text{SO}_4)^2 (\text{H}_2\text{S})^6 (\text{H}^+)^4}$$

but  $(\text{S}^0) = 1$   
 $(\text{H}_2\text{O}) = 1$

$$K = \frac{1}{(\text{SO}_4)^2 (\text{H}_2\text{S})^6 (\text{H}^+)^4} = 10^{43.68} \frac{1}{(\text{SO}_4)^2 (\text{H}_2\text{S})^6 (\text{H}^+)^4} = 10^{21.84}$$

d) Is  $\text{S}^0$  stable with regard to this reaction at the point of reaction in the Black Sea where  $\text{pH} = 7.5$ ,  $\text{SO}_4^{2-} = 17 \text{ mM}$ ,  $\text{H}_2\text{S} = 1 \mu\text{M}$ . Assume for this calculation that concentrations and activities are equal. (10 pts)

$$\begin{aligned} Q &= \frac{(17 \times 10^{-3})^2 (1 \times 10^{-6})^6 (10^{-7.5})^4}{(289 \times 10^{-6}) (10^{-36}) (10^{-30})} \\ &= \frac{1}{(10^{-35.3}) (10^{-36}) (10^{-30})} \\ &= \frac{1}{10^{-69.5}} \end{aligned}$$

$$\begin{aligned} Q &= \frac{1}{(17 \times 10^{-3})^2 (1 \times 10^{-6})^6 (10^{-7.5})^4} \\ Q &= \frac{1}{10^{-34.7}} \\ Q &= 10^{34.7} \\ K &= 10^{21.8} \\ Q/K &> 1 \end{aligned}$$

SO

$$\begin{aligned} Q &= 10^{69.5} \\ K &= 10^{43.7} \\ Q/K &= 10^{25.8} \\ &> 1 \end{aligned}$$

SO REACTION TO LEFT,  $\text{S}^0$  NOT STABLE!

19. Fossil fuel CO<sub>2</sub> is on its way up from the pre-industrial value of about P<sub>CO2</sub> = 300 μatm to about 600 μatm. By one scenario we can assume that alkalinity will stay constant at 2100 μEq.

Remember that the equilibrium constants for CO<sub>2</sub> and carbonic acid are:

$$K_H' = [\text{H}_2\text{CO}_3^*] / P_{\text{CO}_2} = 10^{-1.5}$$

$$K_1' = [\text{HCO}_3^-][\text{H}^+] / [\text{H}_2\text{CO}_3] = 10^{-6.0}$$

$$K_2' = [\text{CO}_3^{2-}][\text{H}^+] / [\text{HCO}_3^-] = 10^{-9.0}$$

$$\left. \begin{array}{l} 300 \times 10^{-6} \\ 3 \times 10^{-4} \\ 10^{-3.52} \end{array} \right\}$$

a) Calculate the [H<sub>2</sub>CO<sub>3</sub>\*] in equilibrium with an atmospheric P<sub>CO2</sub> = 300 μatm. (10 pts)

$$[\text{H}_2\text{CO}_3^*] = K_H P_{\text{CO}_2} = 10^{-1.5} \cdot 10^{-3.52} = 10^{-5.02}$$

$$\left. \begin{array}{l} 600 \times 10^{-6} \\ 6 \times 10^{-4} \\ 10^{-3.22} \end{array} \right\}$$

b) Calculate pH for this open carbonate system when P<sub>CO2</sub> has doubled to 600 μatm. (10 pts)

$$\begin{aligned} \text{Alk} &\approx \text{HCO}_3^- + 2\text{CO}_3^{2-} \\ &= C_T (\alpha_1 + 2\alpha_2) \\ &= \frac{K_H P_{\text{CO}_2}}{\alpha_0} (\alpha_1 + 2\alpha_2) \end{aligned}$$

$$\text{Alk} = K_H P_{\text{CO}_2} \left( \frac{K_1 + 2K_1 K_2}{1 + K_1 + 2K_1 K_2} \right)$$

$$2100 \times 10^{-6} = 10^{-1.5} \cdot 10^{-3.22} \left( \frac{1 \cdot 10^{-6} + 2 \cdot 10^{-6} \cdot 10^{-9}}{1 + 10^{-6} + 2 \cdot 10^{-6} \cdot 10^{-9}} \right)$$

$$2.1 \times 10^{-3} = 10^{-4.72} \left( \frac{1 + 10^{-6} + 2 \times 10^{-15}}{1 + 2} \right)$$

$$10^{-2.67} \cdot 1.2 = 10^{-10.72} \cdot 1.2 - 10^{-19.42}$$

$$10^{-2.67} \cdot 1.2 - 10^{-10.72} \cdot 1.2 + 10^{-19.42} = 0$$

$$1.2 \cdot 10^{-2.67} - 1.2 \cdot 10^{-10.72} + 10^{-19.42} = 0$$

$$\text{pH} \approx 7.97$$

20. Nitrogen fixation has become the latest "hot topic". Lets assess its importance. A Construct a two-box ocean model for fixed nitrogen as nitrate. Include river inflow, atmospheric input to the surface box, nitrification input to the surface box at the air-sea surface, upwelling, downwelling, denitrification in the deep box and a biological flux of organic N (B). The removal rate to the sediments (S) is expressed as  $S = f \times B$ .

From the literature we find so information that we may (or may not) need:

$Q_{DENITRIF}$  Global denitrification = 175 TgN  $y^{-1}$

$f = 0.01$

$Q_R$  River Input = 76 TgN  $y^{-1}$

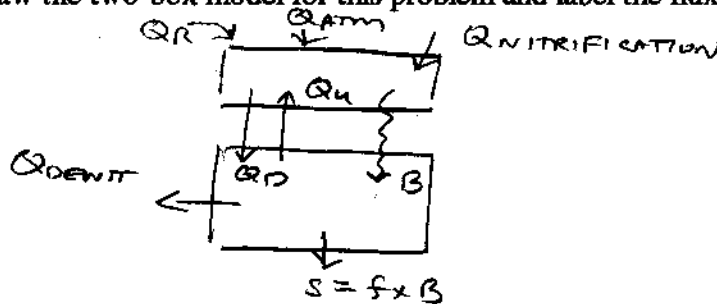
Area =  $3.61 \times 10^8 \text{ km}^2 \times \frac{10^6 \text{ m}^2}{\text{km}^2} \times \frac{10^4 \text{ m}^2}{\text{m}^2}$   
 $= 3.61 \times 10^{18} \text{ cm}^2$

$Q_{ATM}$  Atmospheric Deposition = 30 TgN  $y^{-1}$

$S$  Sediment Burial = 25 TgN  $y^{-1}$

$Q_u, Q_d$  Upwelling rate  $V = 300 \text{ cm } y^{-1}$

a) Draw the two-box model for this problem and label the fluxes. (5 pts)



b) Write the mass balance equations for the surface and deep boxes. (10 pts)

ASSUME Steady STATE

SURFACE  $0 = Q_R + Q_{NITRIF} + V_m C_D - V_m C_S - B$

DEEP  $0 = V_m C_S - Q_{DENITRIF} + B - fB - V_m C_D$

c) If the average concentration of nitrate in the deep box is  $C_D = 30 \mu\text{M}$ , and the surface box is  $C_S = 0 \mu\text{M}$ , what is the rate of nitrogen fixation required for a steady state balance? (10 pts)

SURF  $0 = Q_R + Q_{NITRIF} + V_m C_D - B \quad \therefore B = Q_R + Q_{NITRIF} + V_m C_D$

DEEP  $0 = B - Q_{DENITRIF} - f \times B - V_m C_D$

$= (1-f)B - Q_{DENITRIF} - V_m C_D$

$= (1-f)(Q_R + Q_{NITRIF} + V_m C_D) - Q_{DENITRIF} - V_m C_D$

$0 = (1-f)Q_R + (1-f)Q_{NITRIF} + (1-f)V_m C_D - Q_{DENITRIF} - V_m C_D$

$(1-f)Q_{NITRIF} = Q_{DENITRIF} + V_m C_D - (1-f)V_m C_D - (1-f)Q_R$

$Q_{NITRIF} = 175 + 0.01(300 \text{ cm } y^{-1}) 30 \times 10^{-6} \text{ mol } / 10^3 \text{ m}^3 \times 3.61 \times 10^{18} \text{ cm}^2 - (0.99)(76)$

d) What is its importance relative to river, atmospheric and upwelling input? (5 pts)

$= 175 + 4.5 - 76 \text{ Tg } y^{-1}$

$Q_{NITRIF} = 104 \text{ Tg } y^{-1}$

$Q_{NITRIF}$  larger than river & much larger than upwelling  
 larger than atmospheric input

